

# River channel formation and response to variations in discharge, sediment and vegetation

## PART 1: RIVER CHANNEL GEOMETRY

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Tagliamento River, Italy



XXXVIII

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# PART 1 RIVER CHANNEL GEOMETRY

## PART 1

1. Alluvial rivers
2. Factors governing the channel slope and depth
3. Comparison with empirical relations
4. Factors governing the channel width
5. Effects of discharge, sediment, vegetation

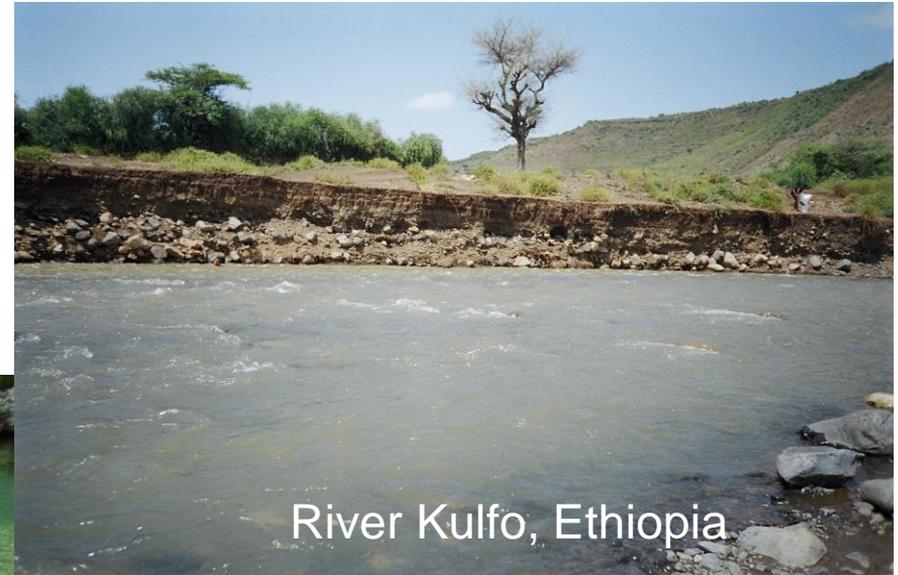
# 1 Alluvial rivers

River channels are excavated by water flowing on erodible soils



River Cellina, Italy

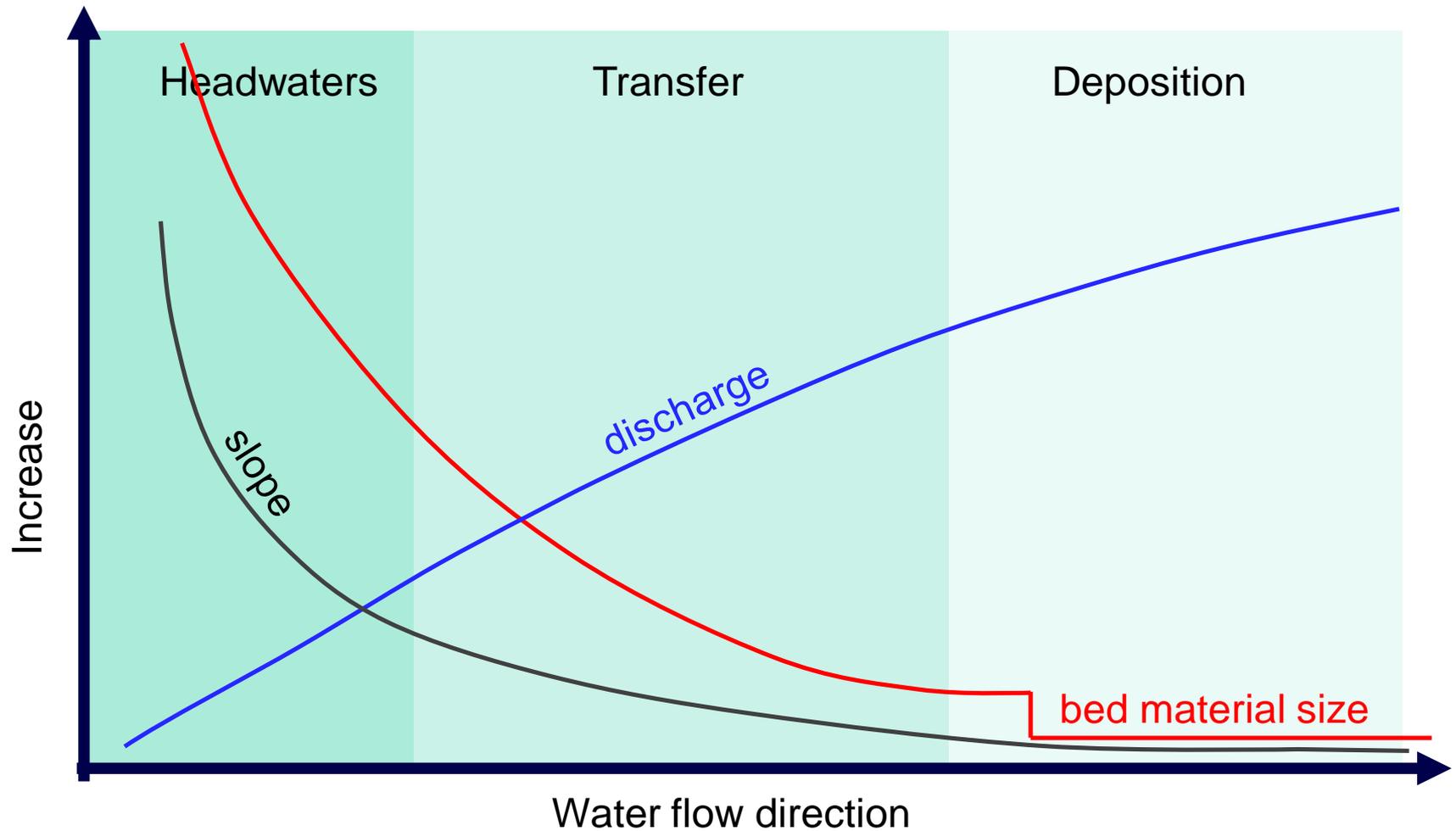
# Definition of alluvial river



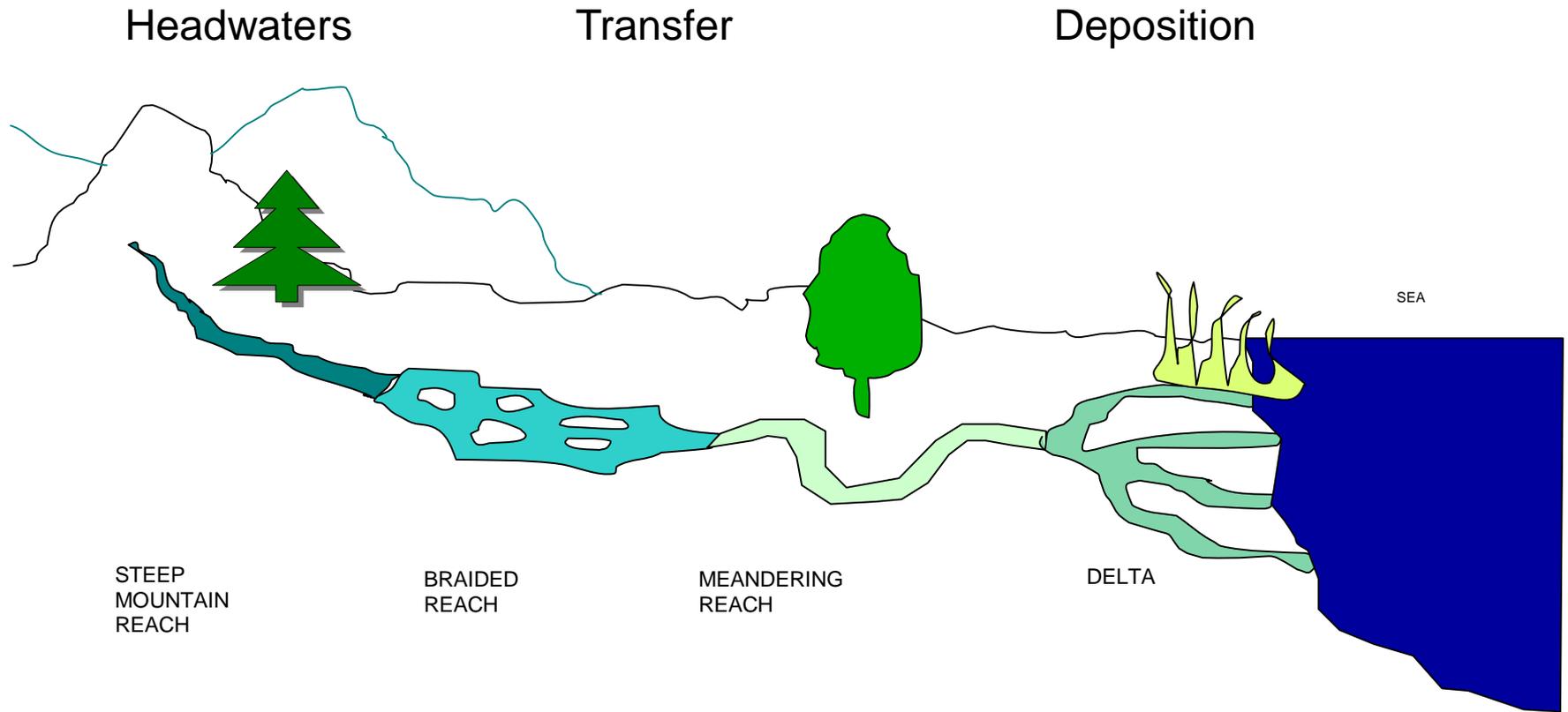
Alluvial river:  
bed made of sediment captured  
and transported by water flow

Erosion and accretion are the  
main processes that shape  
alluvial rivers

# Discharge, slope, sediment



# River planform



# Mountain reaches

River Soča, Slovenia



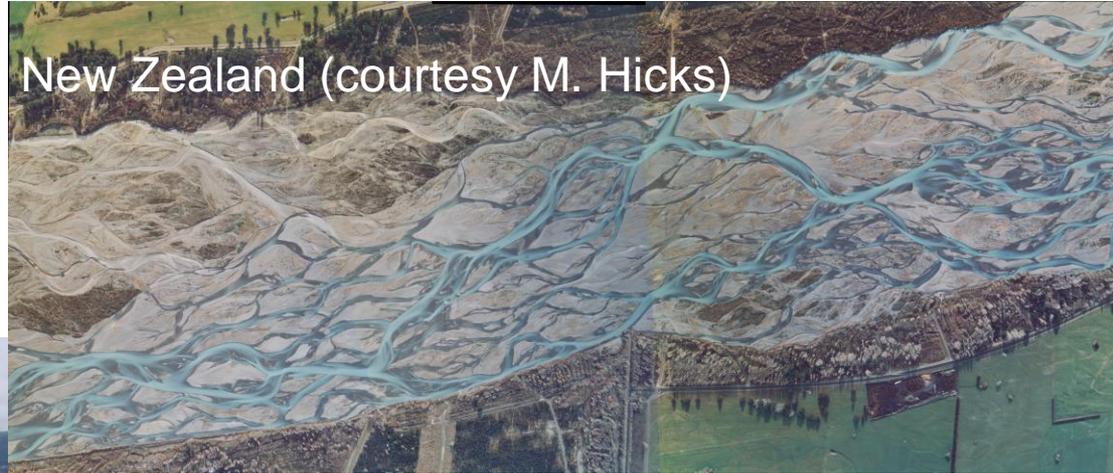
River Mow, Bhutan



Steep, laterally limited by rocks,  
flash floods, large sediment sizes.

# Braided reaches

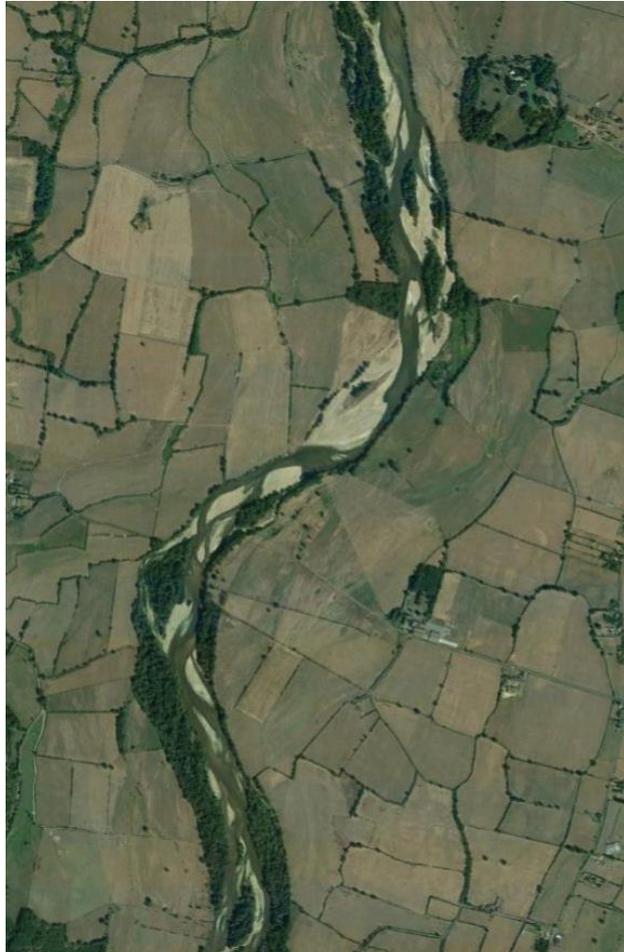
Waimakariri River New Zealand (courtesy M. Hicks)



Tagliamento River, Italy

**TYPICAL OF PIED-MONT AREAS**  
Channels separated by multiple bars emerging during low flow. Flash floods, gravel bed, large sediment inputs. Dominated by erosive processes.

# Transition reaches



River Allier, France (Google Earth image)

McLeod River, Canada (courtesy J. Slaney)



TYPICAL OF PIEDMONT AREAS AND VALLEYS  
Central bars and side channels.  
Rather sinuous course.

# Meandering reaches

River Geul, the Netherlands



River Mara, Tanzania



**TYPICAL OF WIDE VALLEYS**  
Sand-bed and cohesive banks.  
Balanced bank accretion and erosion leading to transverse channel shift and bend growth.

# Anabranching reaches

Zambezi River, Zambia (courtesy E. Mwelwa Mutekenya)



Amazon River, Peru (courtesy E. Mosselman)

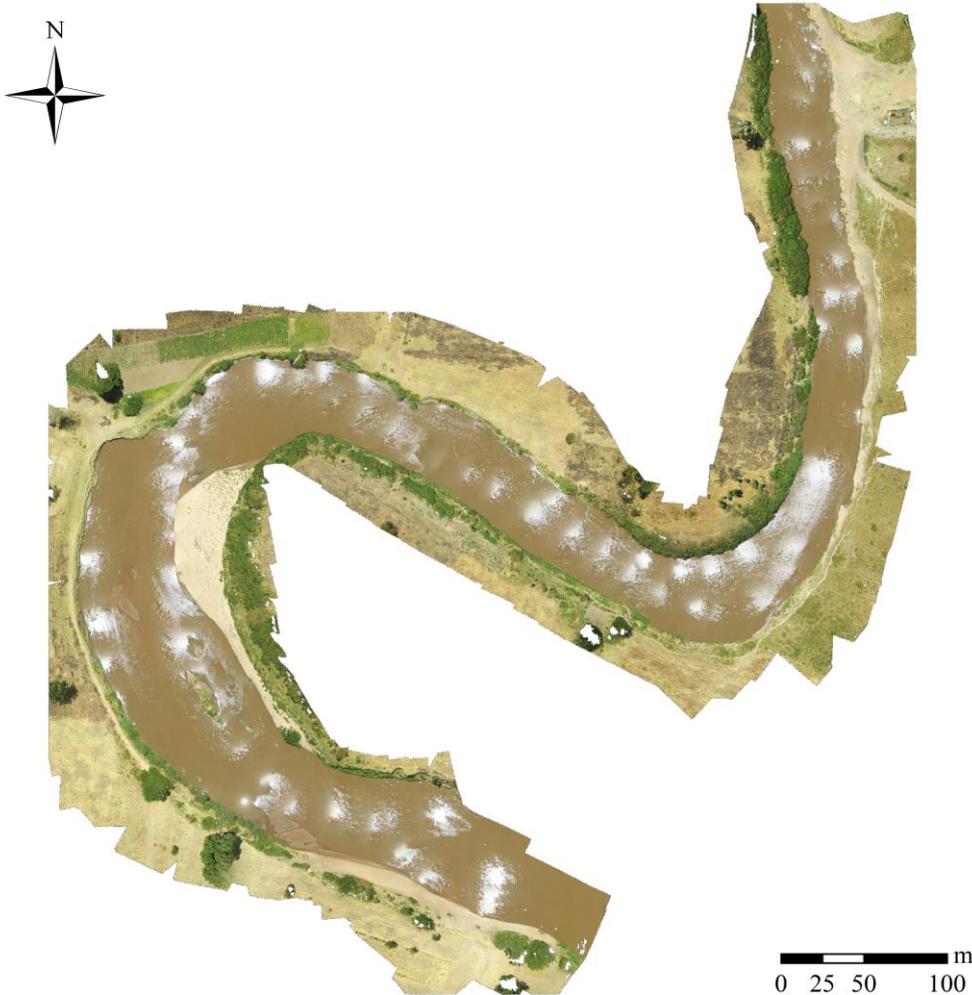


TYPICAL OF LARGE RIVERS IN  
WIDE VALLEYS  
Multiple channels separated by wide  
vegetated islands.

Why do rivers change their slope, size and shape from the mountains to the sea?



# Focus on river reach characteristics

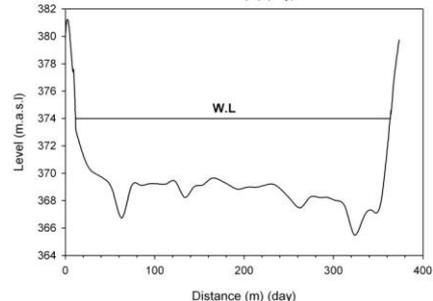
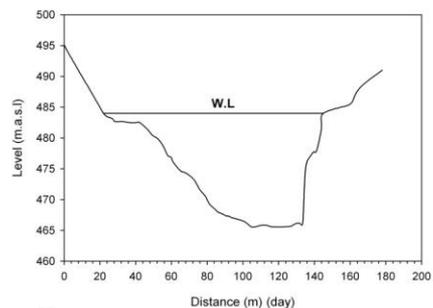
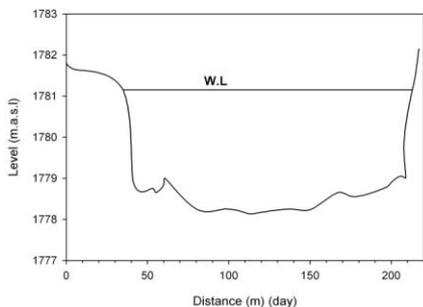


River reach characteristics:

- longitudinal bed slope
- average width
- average depth
- planform

Mara River, Tanzania  
(courtesy F. Bregoli)

# Attention also to cross-section characteristics



Blue Nile, Sudan (courtesy Y. Ali)

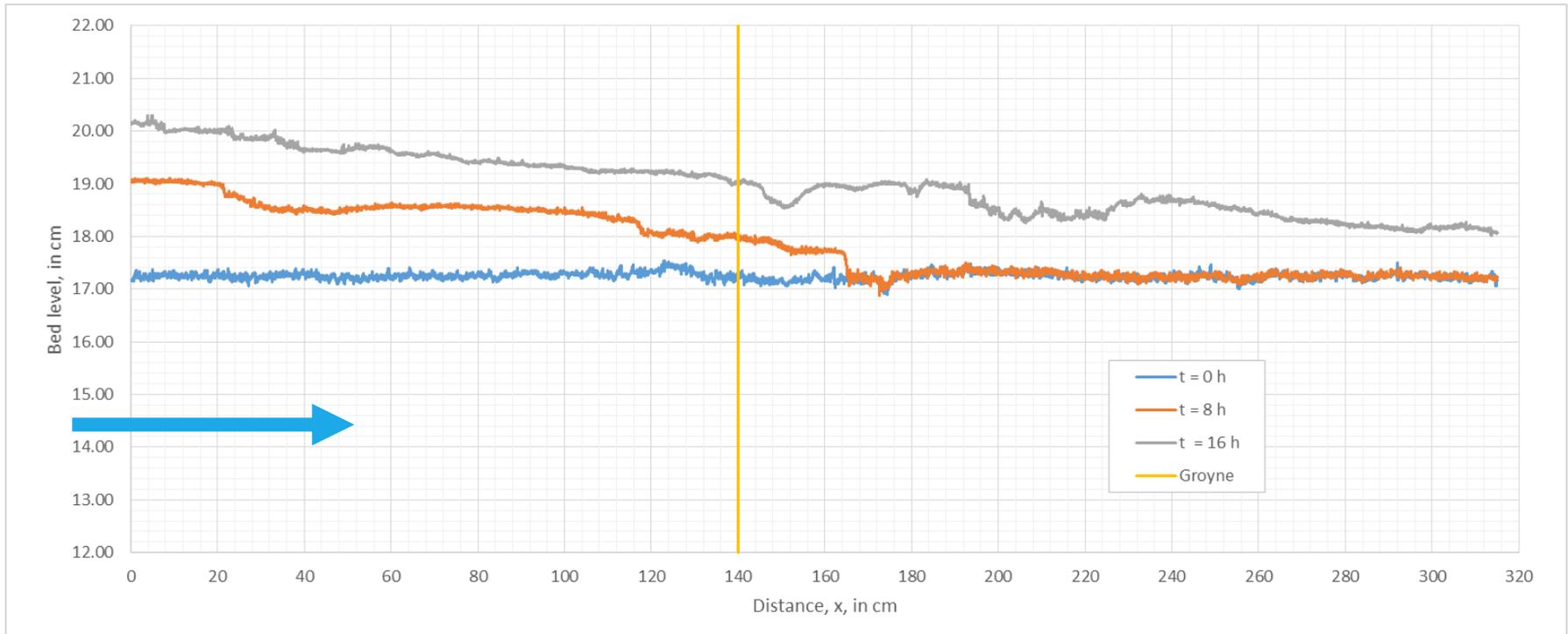
Characteristics of river cross-sections: width, depth, bed topography (bars)



River Mara, Tanzani (courtesy F. Bregoli)

# 2 Factors governing the channel slope and depth

(flume experiments by Crosato, Bonilla-Porrás et al. 2018)

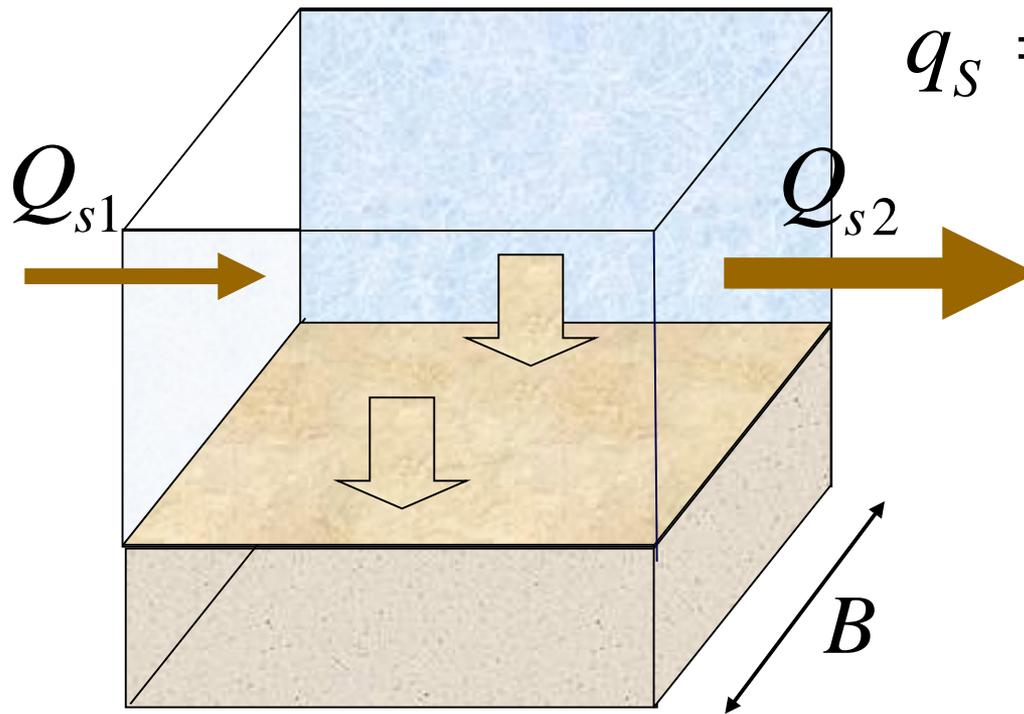


Slope adaptation to boundary conditions by sedimentation front propagation  
Different slope is obtained with different roughness

Morphological changes occur if  
sediment input  $\neq$  sediment output

$$Q_s = Bq_s$$

$$q_s = M(u - u_c)^b$$



$$Q_{s1} = Q_{s2} \text{ if } u_1 = u_2$$

(Exner approach)

# Morphodynamic equilibrium: uniform sediment transport ↔ uniform flow

$$1) u = C\sqrt{hi}$$

Chézy equation for large and shallow rivers:  
momentum equation

$$2) Q_W = uBh$$

continuity equation for water

$$3) q_S = M(u - u_c)^b$$

sediment transport capacity formula (replaces  
momentum equation for sediment)

$$4) Q_S = Bq_S$$

continuity equation for sediment (sediment  
balance)

## ASSUMPTIONS:

$Q_W$  = formative discharge

$u_c \approx 0$  (sand)

Combination of 1 to 4 leads to:

channel depth

$$h = \frac{M^{1/b} Q_W}{Q_S^{1/b} B^{(1-1/b)}}$$

channel slope

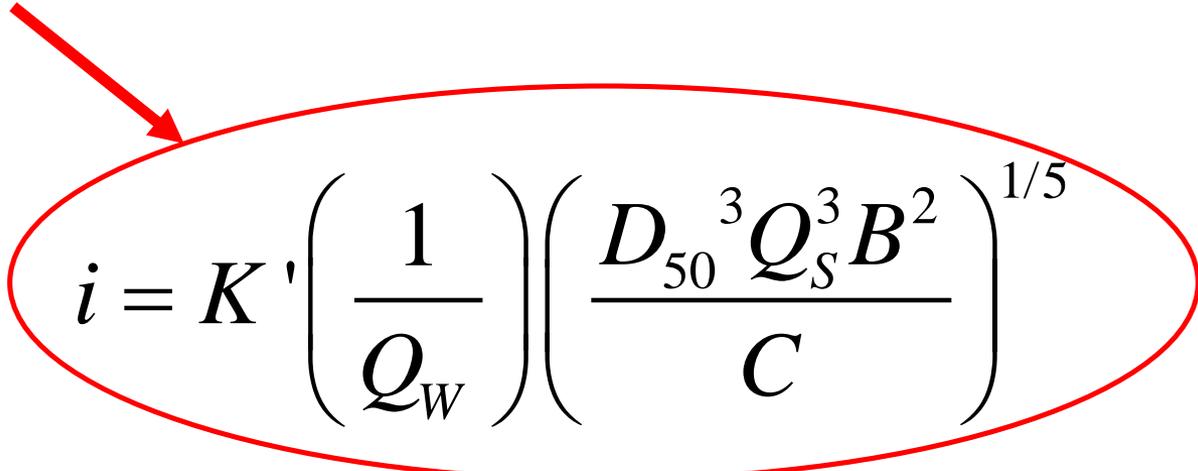
$$i = \frac{Q_S^{3/b} B^{(1-3/b)}}{M^{3/b} Q_W C^2}$$

(morphodynamic equilibrium)

According to Engelund & Hansen (1967)  $b = 5$ ,  $u_c = 0$   
 (sand-bed rivers) and

$$M = K \frac{1}{C^3 \Delta^2 D_{50} \sqrt{g}}$$

$$i = \frac{Q_S^{3/b} B^{(1-3/b)}}{M^{3/b} Q_W C^2}$$



$$i = K' \left( \frac{1}{Q_W} \right) \left( \frac{D_{50}^3 Q_S^3 B^2}{C} \right)^{1/5}$$

(with  $K' = \text{constant}$ )

# Response of channel slope to variations of..

The channel slope  $i$  increases with

$$i \sim \left( \frac{1}{Q_W} \right) \left( \frac{D_{50}^3 Q_S^3 B^2}{C} \right)^{1/5}$$

- Sediment transport:  $Q_S$
- Sediment size:  $D_{50}$
- Channel width:  $B$
- Bed roughness:  $1/C$

The channel slope  $i$  decreases with

- Discharge  $Q_W$

# Response of channel depth to variations of..

The channel depth  $h$  increases with

$$h \sim (Q_W) \left( \frac{1}{B^4 C^3 D_{50} Q_S} \right)^{1/5}$$

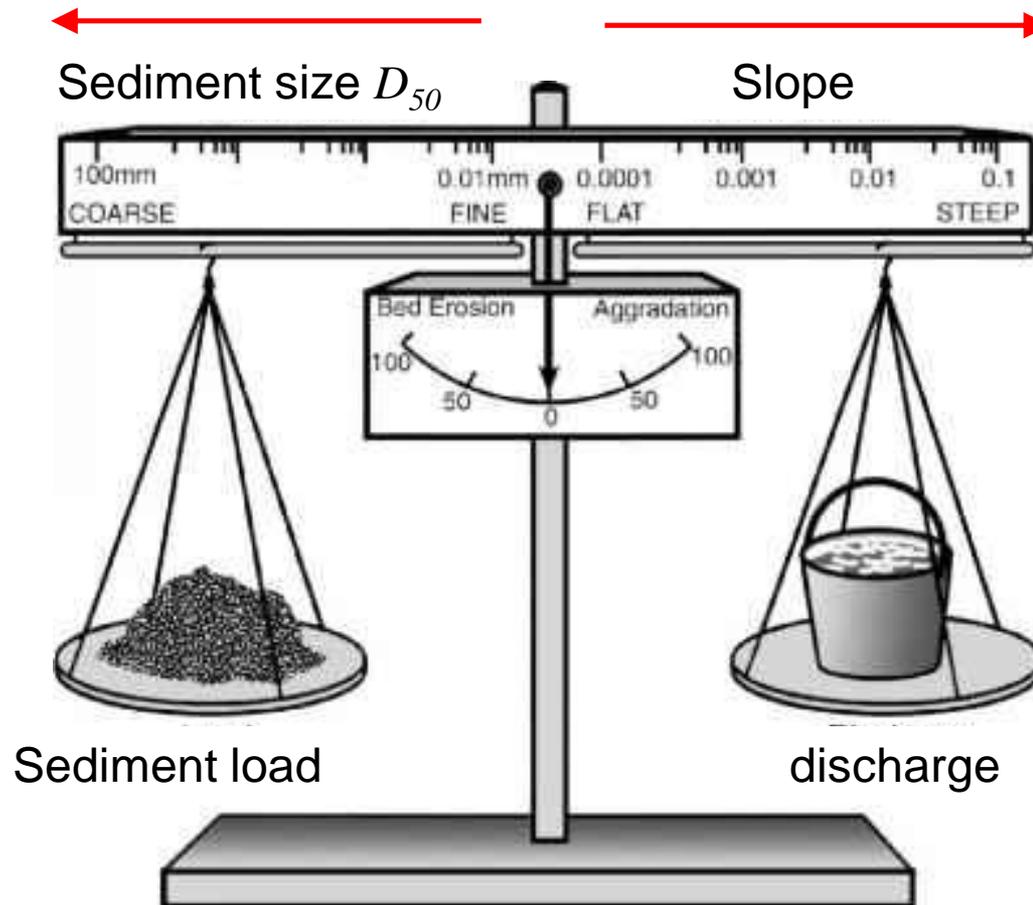
- Discharge  $Q_W$
- Bed roughness  $1/C$

The channel depth  $h$  decreases with

- Channel width:  $B$
- Sediment size:  $D_{50}$
- Sediment load:  $Q_S$

# 3 Comparison with empirical relations

(Lane's balance adapted from Brierley and Fryirs 2005)



Lane's (1955) balance:

Leopold and Maddock  
(1953)

Parker et al. (2007)

$$Q_S D_{50} \sim i Q_W$$

$$i = \alpha_i \left( \frac{1}{Q_{bf}} \right)^{\beta_i}$$

$$i = 0.101 \left( \frac{D_{50}^2 \sqrt{g \Delta D_{50}}}{Q_{bf}} \right)^{0.344}$$

$$i \sim \frac{Q_S D_{50}}{Q_W}$$

$Q_{bf}$  (bankfull discharge)  
assumed as formative

Derived equilibrium law

$$i \sim \left( \frac{1}{Q_W} \right) \left( \frac{D_{50}^3 Q_S^3 B^2}{C} \right)^{1/5}$$

Major difference:  
includes width, assumed as  
known, and bed roughness

# Factors governing the channel slope and depth

Discharge

Sediment size  
Sediment load  
Bed roughness

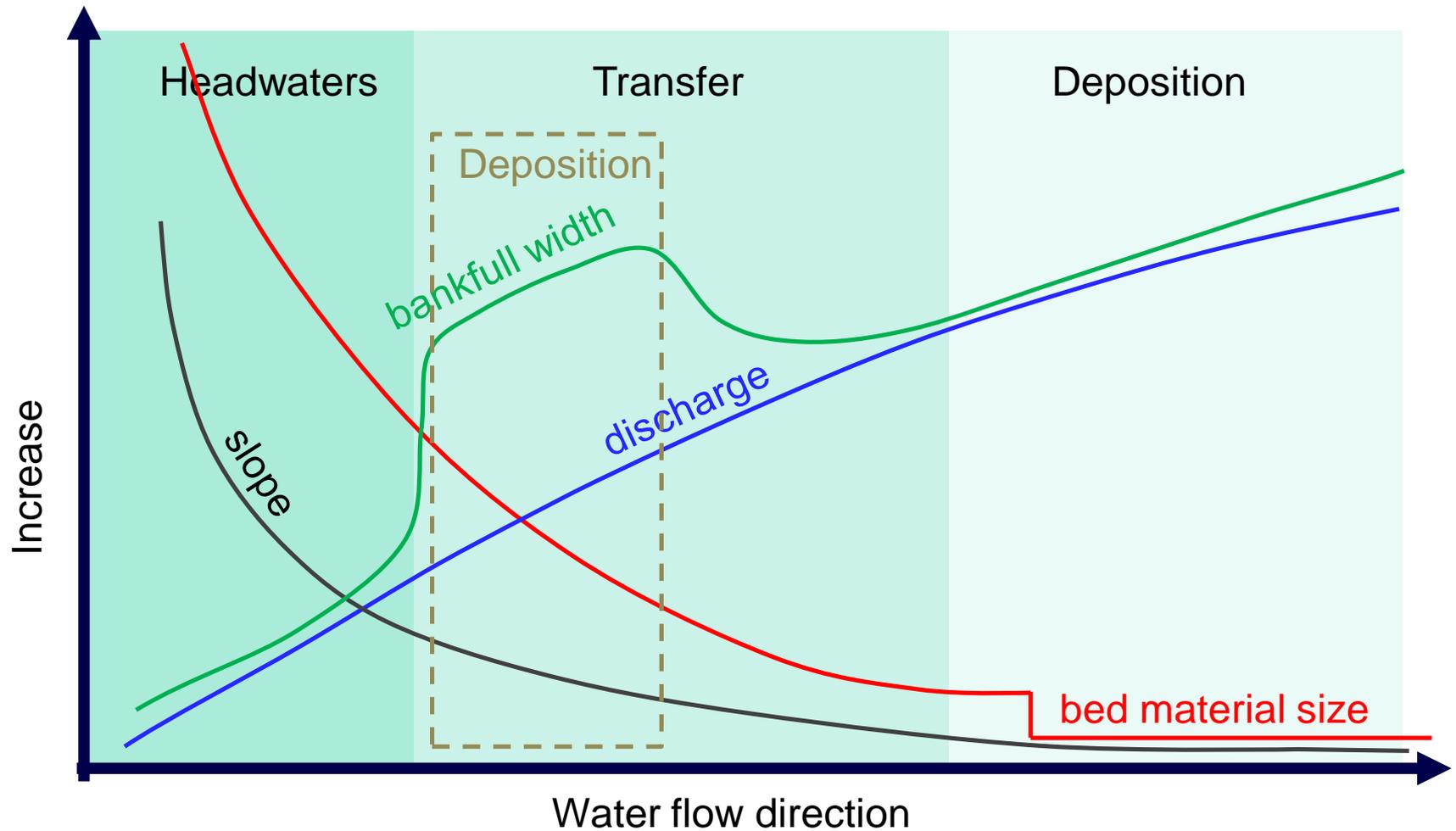
Sediment

Channel width

And what about the channel width?



# Discharge, slope, sediment and width

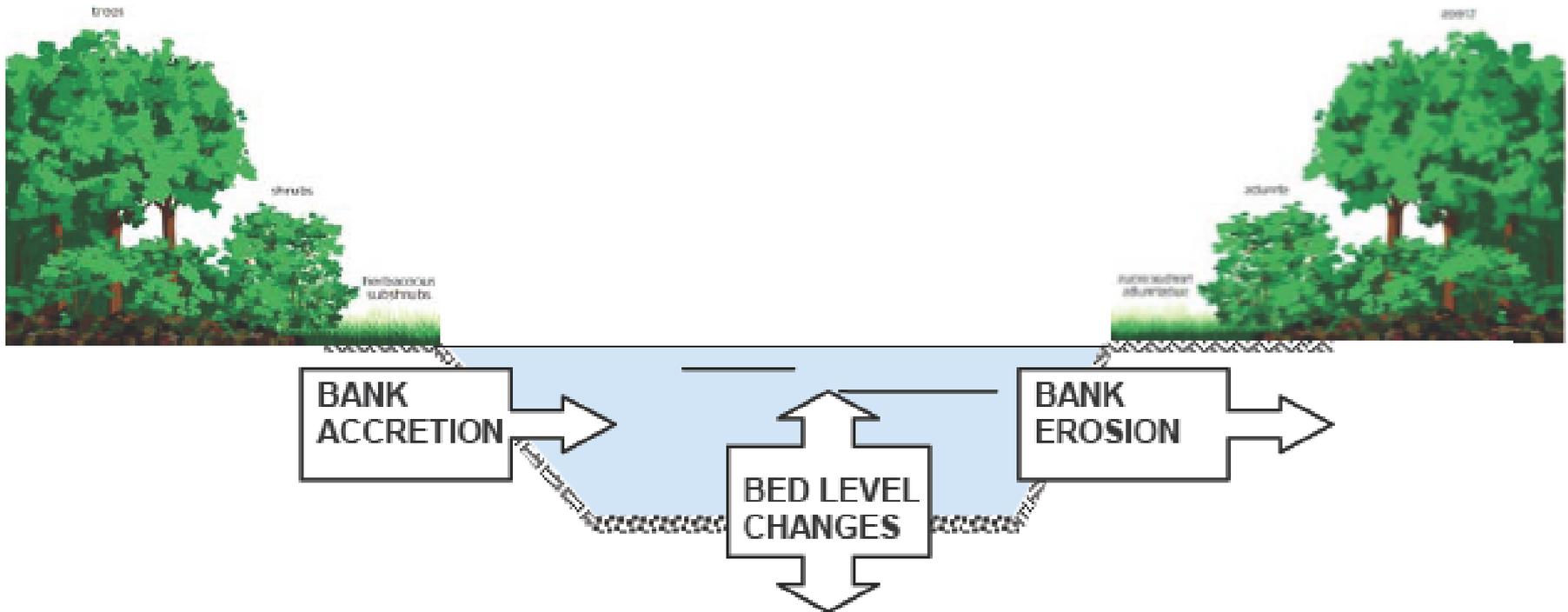


# 4 Factors governing the channel width



River Pilcomayo, Paraguay

# River width processes: bank erosion and accretion



# Bank erosion

1. Flow capturing sediment particles directly from the bank
2. Bank failure caused by geotechnical instability, triggered by bank toe erosion



# Capturing of sediment particles dominates non-cohesive banks



River Cellina, Italy

# Failure dominates steep cohesive banks



River Meuse, the Netherlands (courtesy G. Duró)

# Bank accretion

1. Near-bank sediment deposition
2. Soil stabilization by plants and consolidation processes
3. Sedimentation and level raise



Colorado River, Colorado



# Vegetation affects bank erosion and accretion by:

(bio-stabilization)

- Reinforcing soil by roots
- Protecting soil by cover
- Decreasing local flow velocity
- Deflecting the flow

- Colonization
- Enhancing local deposition
- Favoring vegetation growth

(bio-engineering)

Sarca River, Italy

# Factors governing the channel width

Bank material (bank erodibility)

Sediment (near-bank deposition/scour)

Discharge (sediment entrainment/scour)

**FOCUS**

Vegetation (bank erodibility, bank accretion)

Groundwater flow (bank erodibility)

# Existing empirical relations for channel width

Example: Leopold and Maddock (1953)

$$B = \alpha_B Q_W^{\beta_B}$$

power law of  
“bankfull discharge”

Channel width and depth as a function of discharge and sediment size:

- Parker et al. (2007): gravel-bed rivers
- Wilkerson and Parker (2011): sand-bed rivers

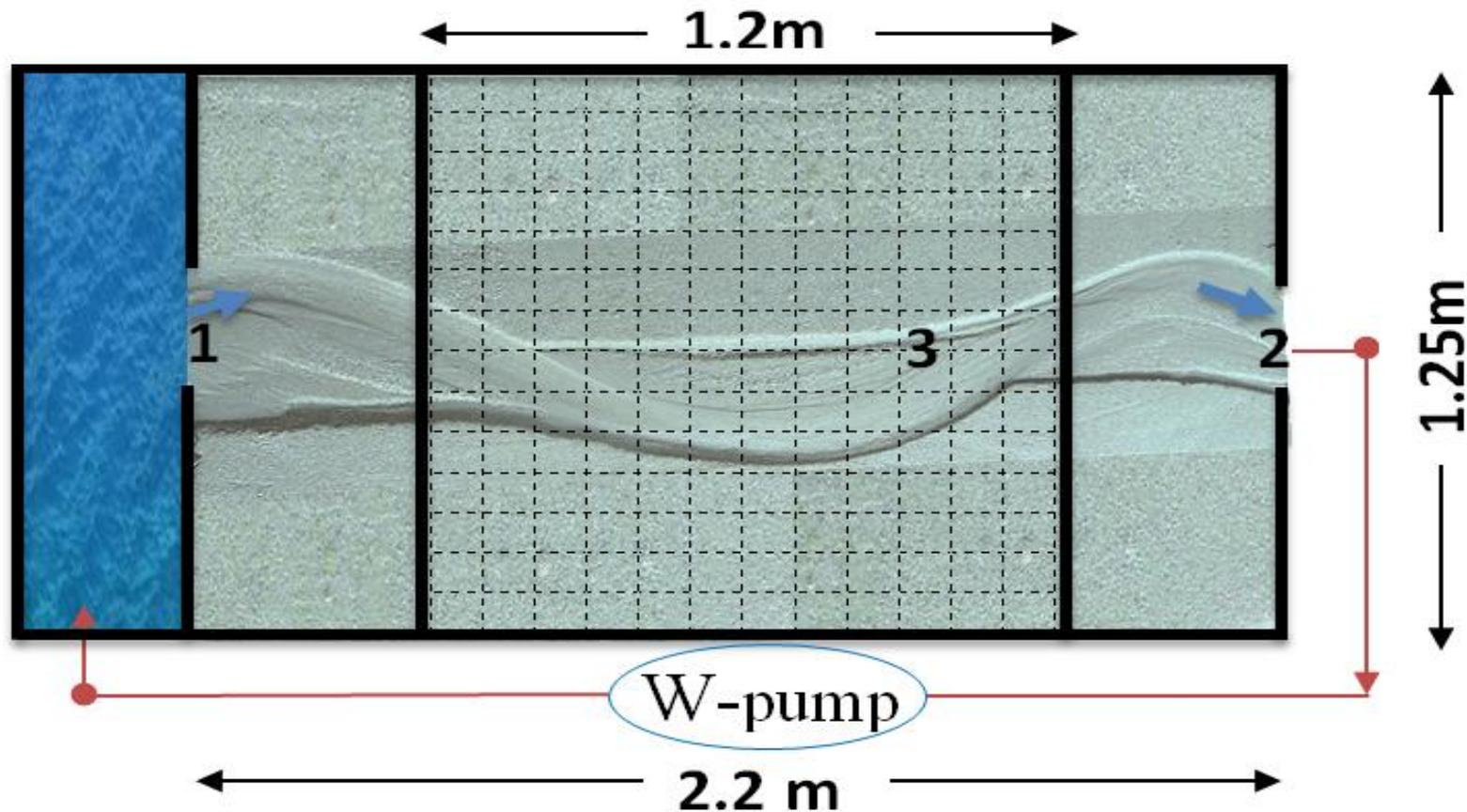
among others

# 5 Effects of discharge, sediment, vegetation

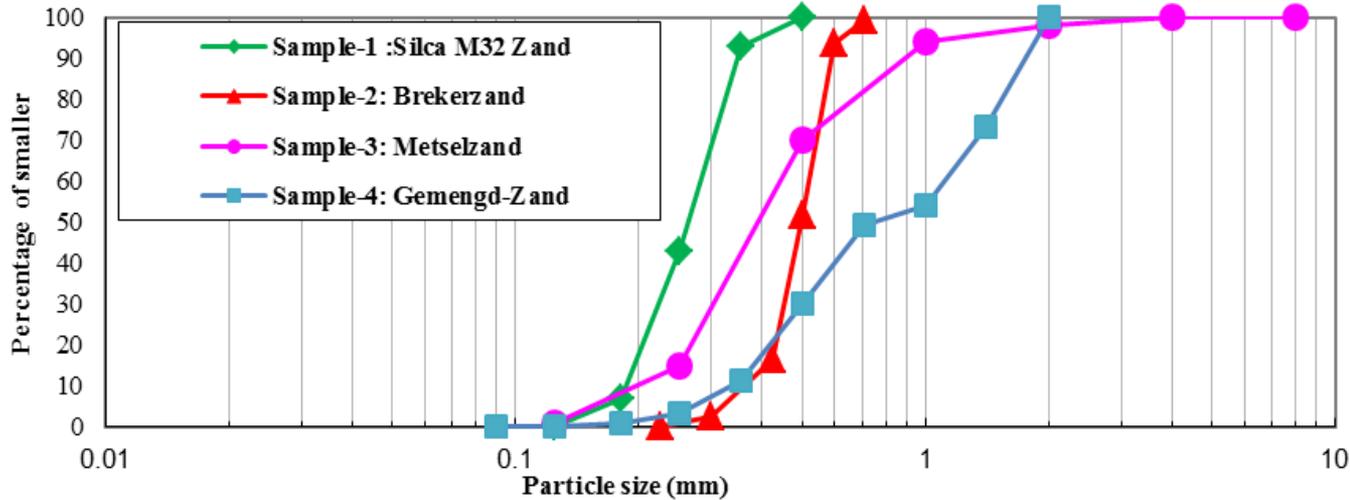
Results of some recent studies

# Effects of discharge variability and sediment on channel width - Experimental study

(Byishimo, 2014; Vargas-Luna et al., 2018)



## Granulometric curves of Sediment samples



Sample 1  $D_{50} = 0.26 \text{ mm}$



Sample 2  $D_{50} = 0.5 \text{ mm}$



Sample 3  $D_{50} = 0.4 \text{ mm}$



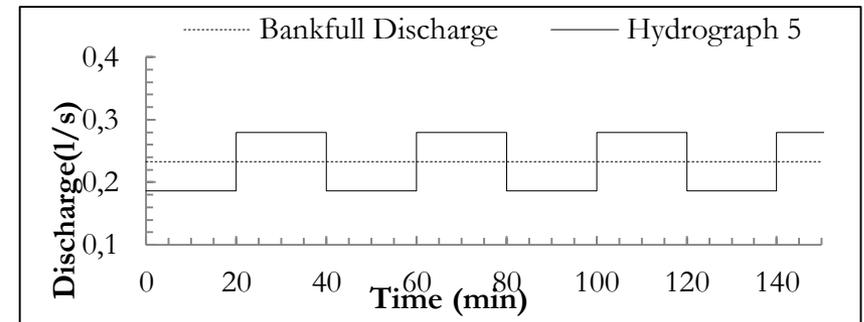
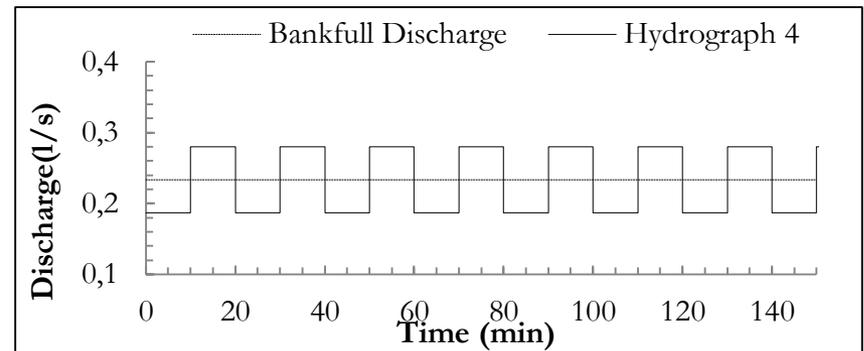
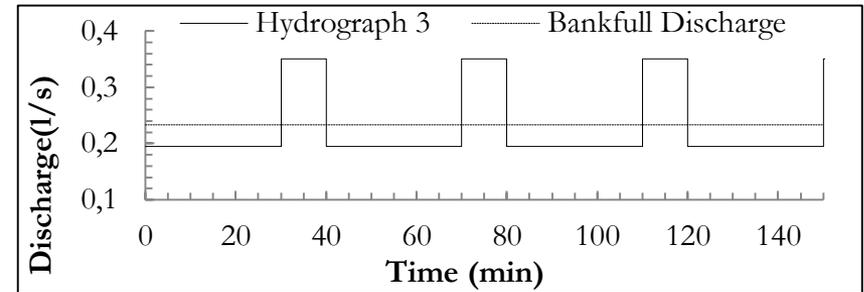
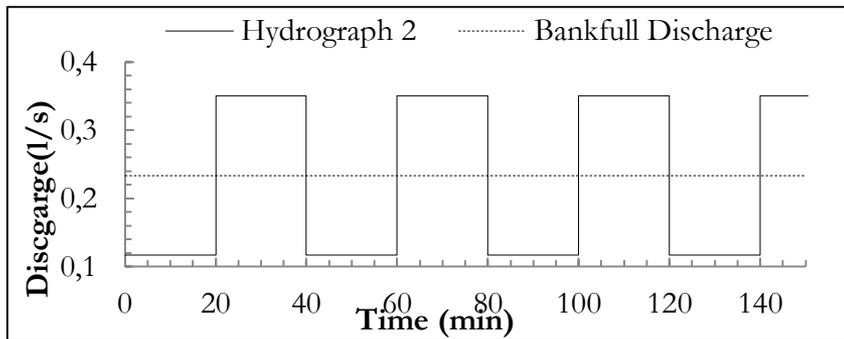
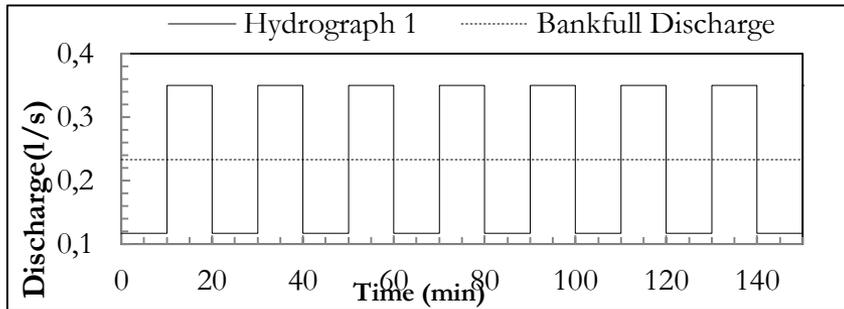
Sample 4  $D_{50} = 0.7 \text{ mm}$



Sample	$D_{50}$ (mm)	Sorting index ( $I$ )
1	0.26	1.34
2	0.50	1.23
3	0.40	1.80
4	0.70	2.16

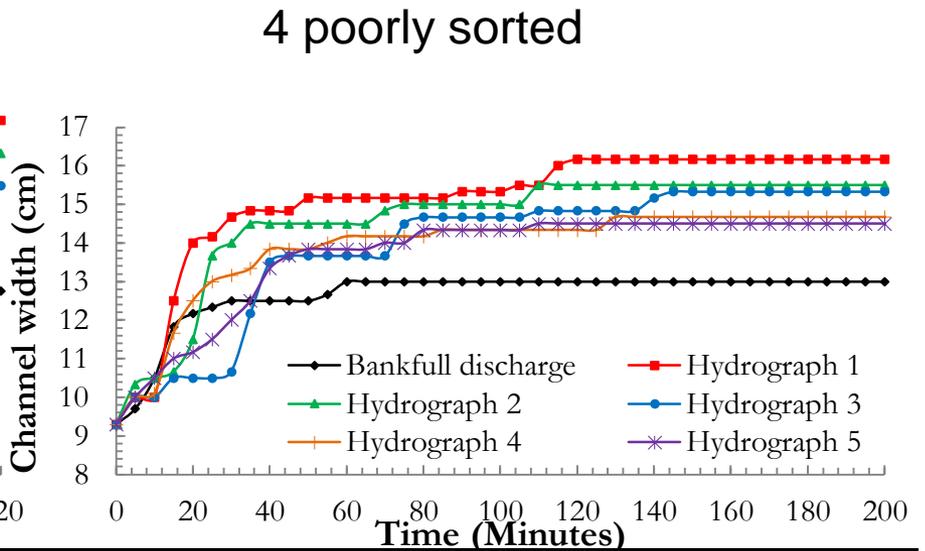
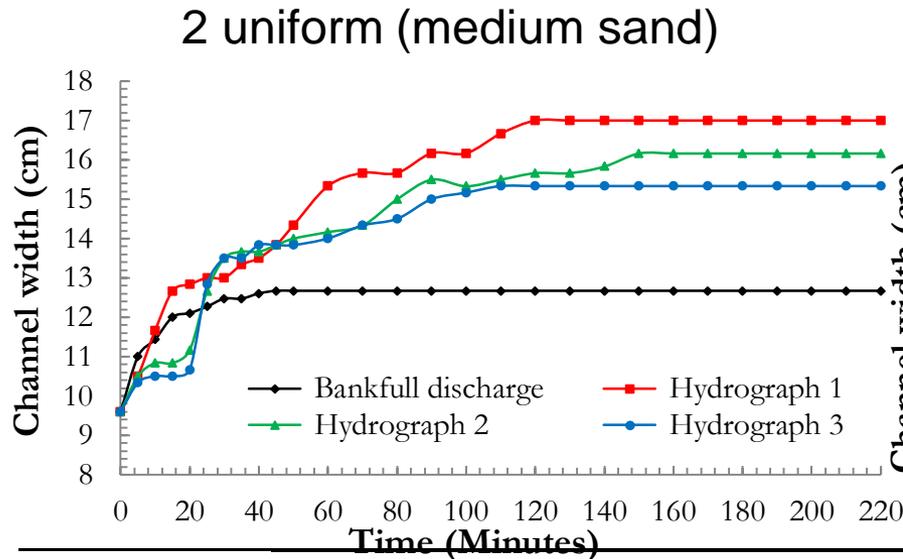
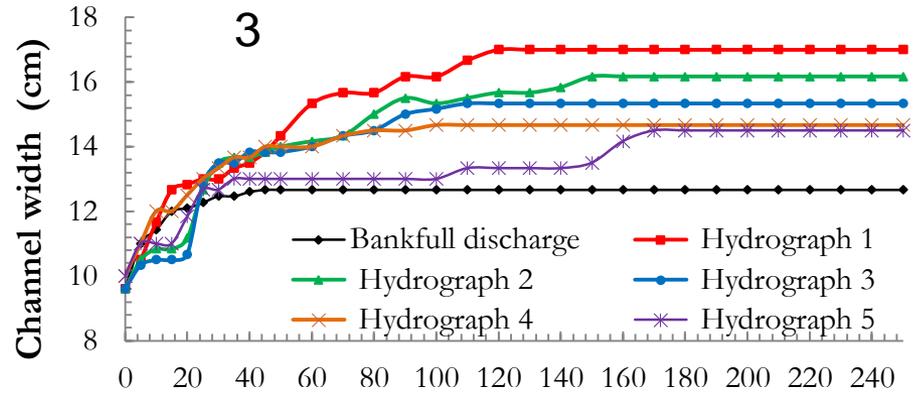
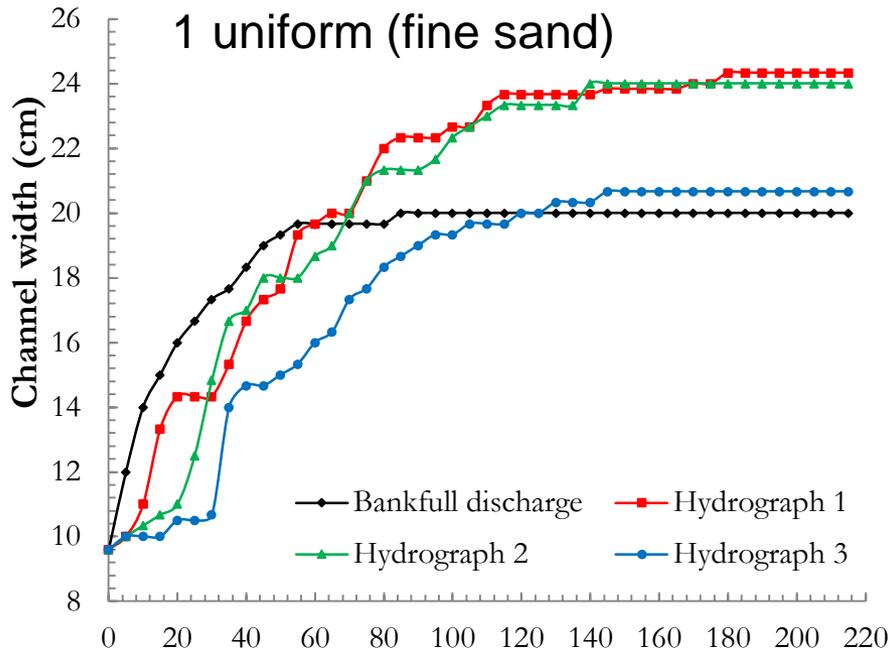
$$I = 0.5 \left( \frac{D_{16}}{D_{50}} + \frac{D_{84}}{D_{50}} \right)$$

# 6 discharge regimes all with the same average



Constant (average) discharge

# Results discharge & sediment



# Results: effects of sediment (constant discharge)

Sample 1, uniform fine sand



Sample 4, poorly sorted sand

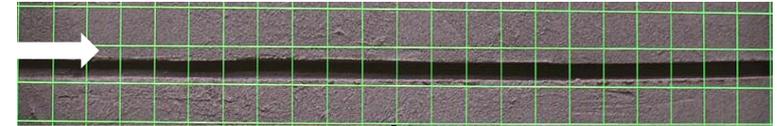


(at the scale of the flume poorly sorted sediment behaves as cohesive)

# Results formative discharge

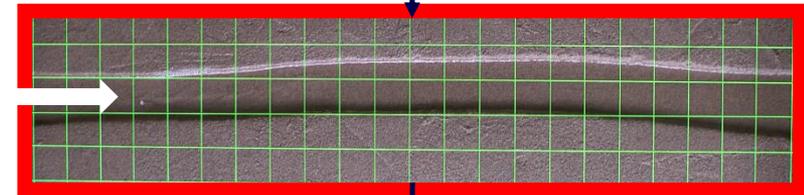
The geometric bankfull discharge is not the formative discharge

start

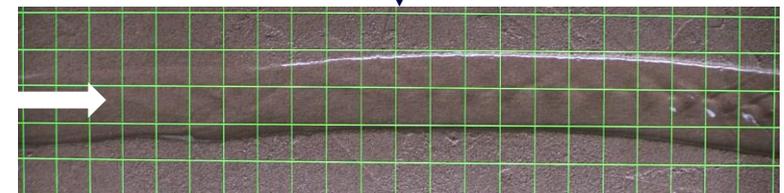


channel evolution with constant (formative) discharge

geometric bankfull discharge assessment



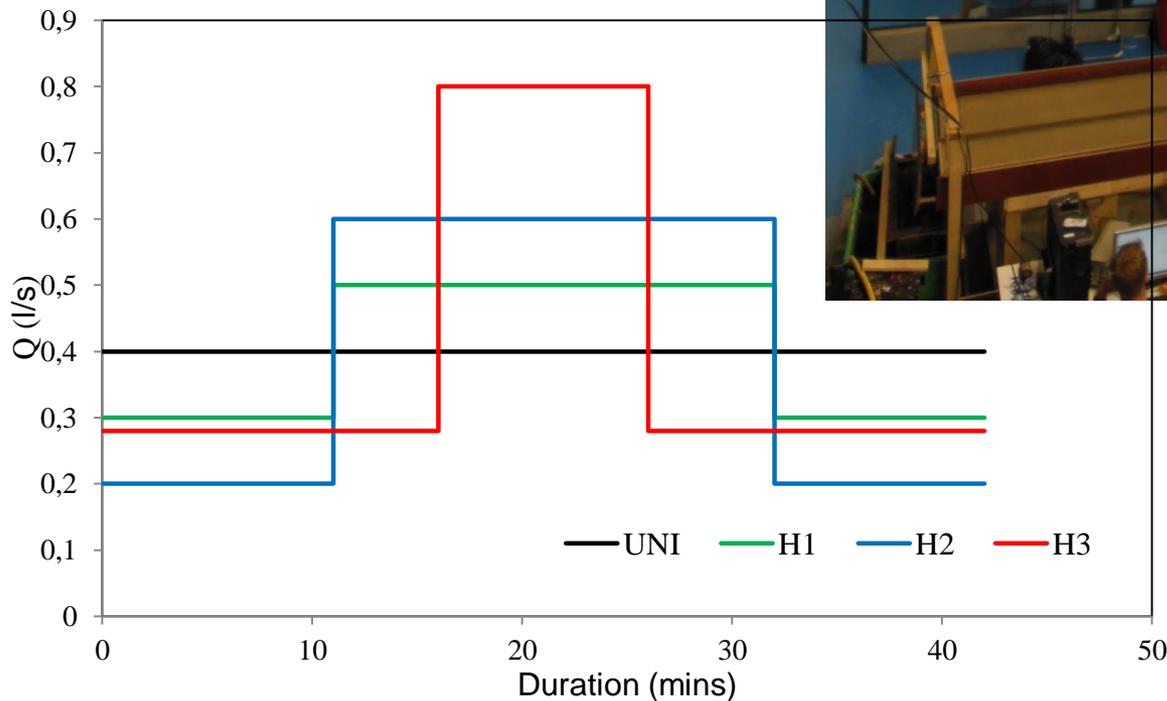
further channel evolution with assessed geometric bankfull discharge



With variable flow the formative discharge corresponds to the frequent peak flow  
From literature: in real rivers, the flow with return period of 1 to 2 years

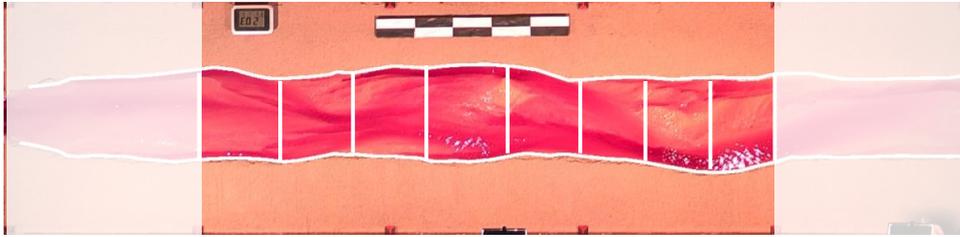
# Effects of starting condition and sediment load on channel width – Experimental study

(Singh, 2015)

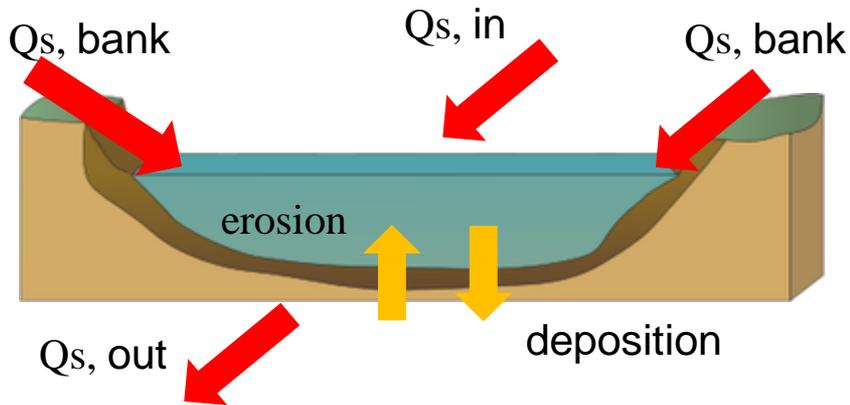


Variable discharge  
4 different starting widths  
Sediment:  $D_{50} = 1$  mm  
(poorly sorted)

## Channel width measurement



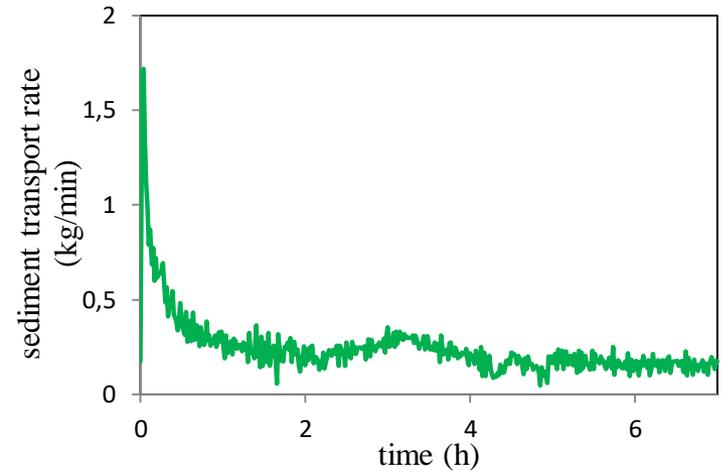
## Sediment budget calculations



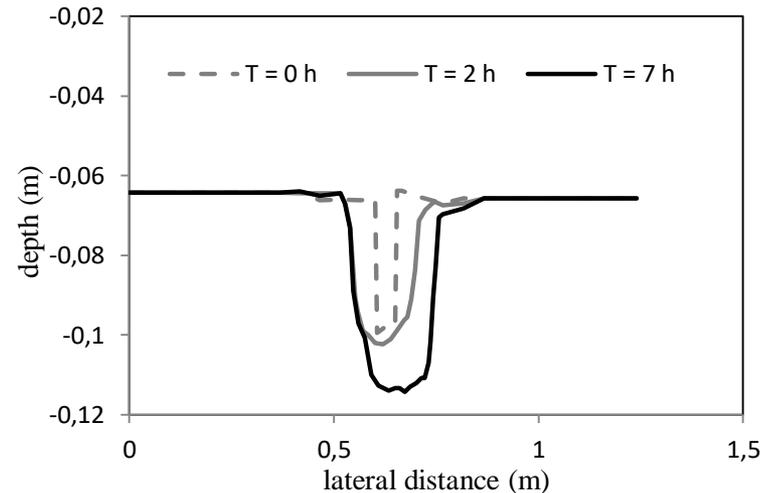
Four different initial widths, 0.04 m, 0.1 m, 0.25 m and 0.4 m

- Without sediment supply
- With sediment supply:  $Q_s = 90$  (g/min)

## Sediment transport rate measurement



## Cross-section measurement



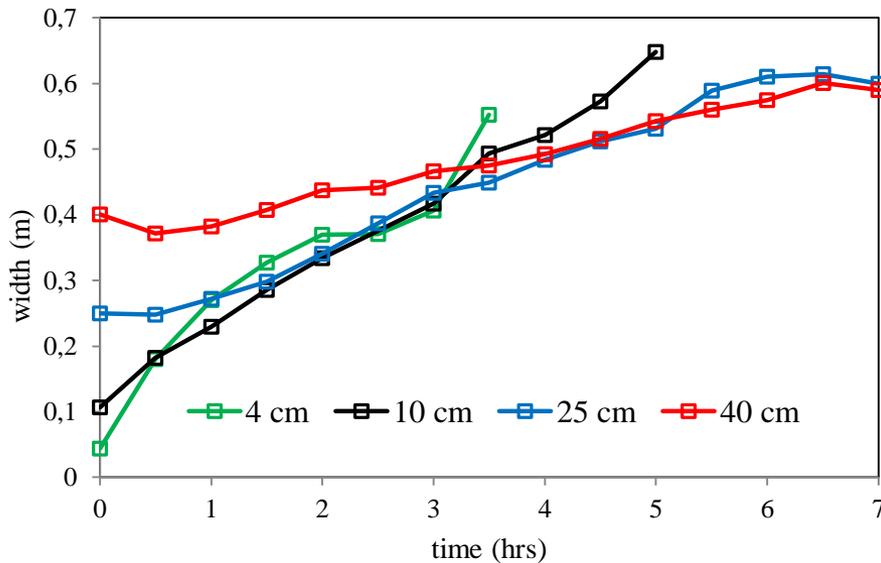
# Results: effects of starting $B$ and sediment load

Constant discharge

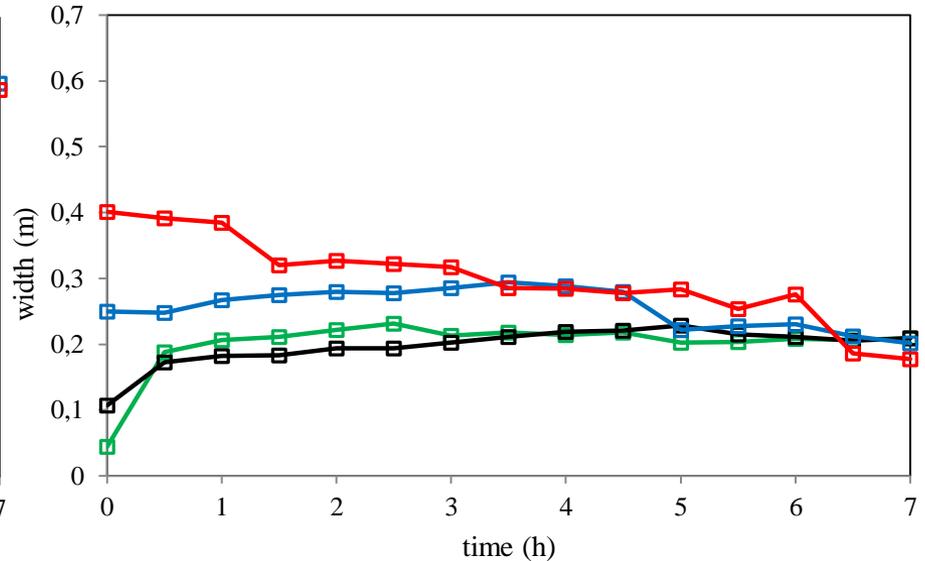
with upstream sediment supply

no upstream sediment supply

Width evolution in time



Width evolution in time



Narrower initial channels => wider channels

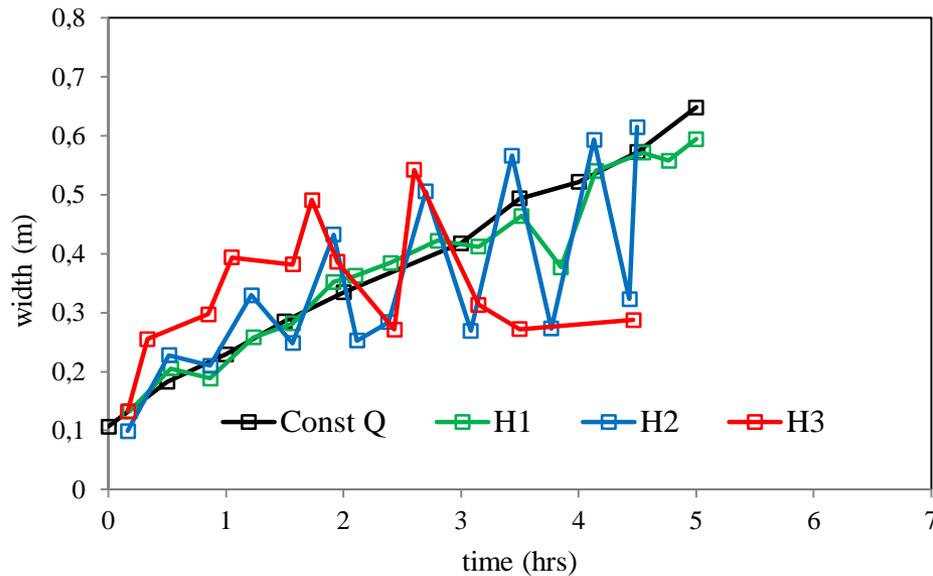
Same final river width

# Results: effects of discharge and sediment load

## Different discharge regimes

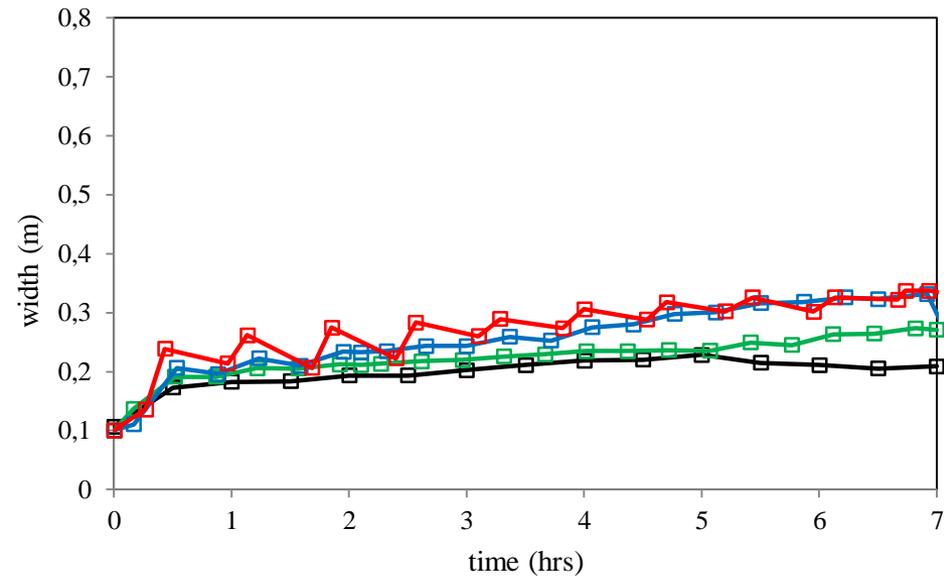
with upstream sediment supply

Width evolution in time



no upstream sediment supply

Width evolution in time



Note: same sediment supply rate of 90 g/min

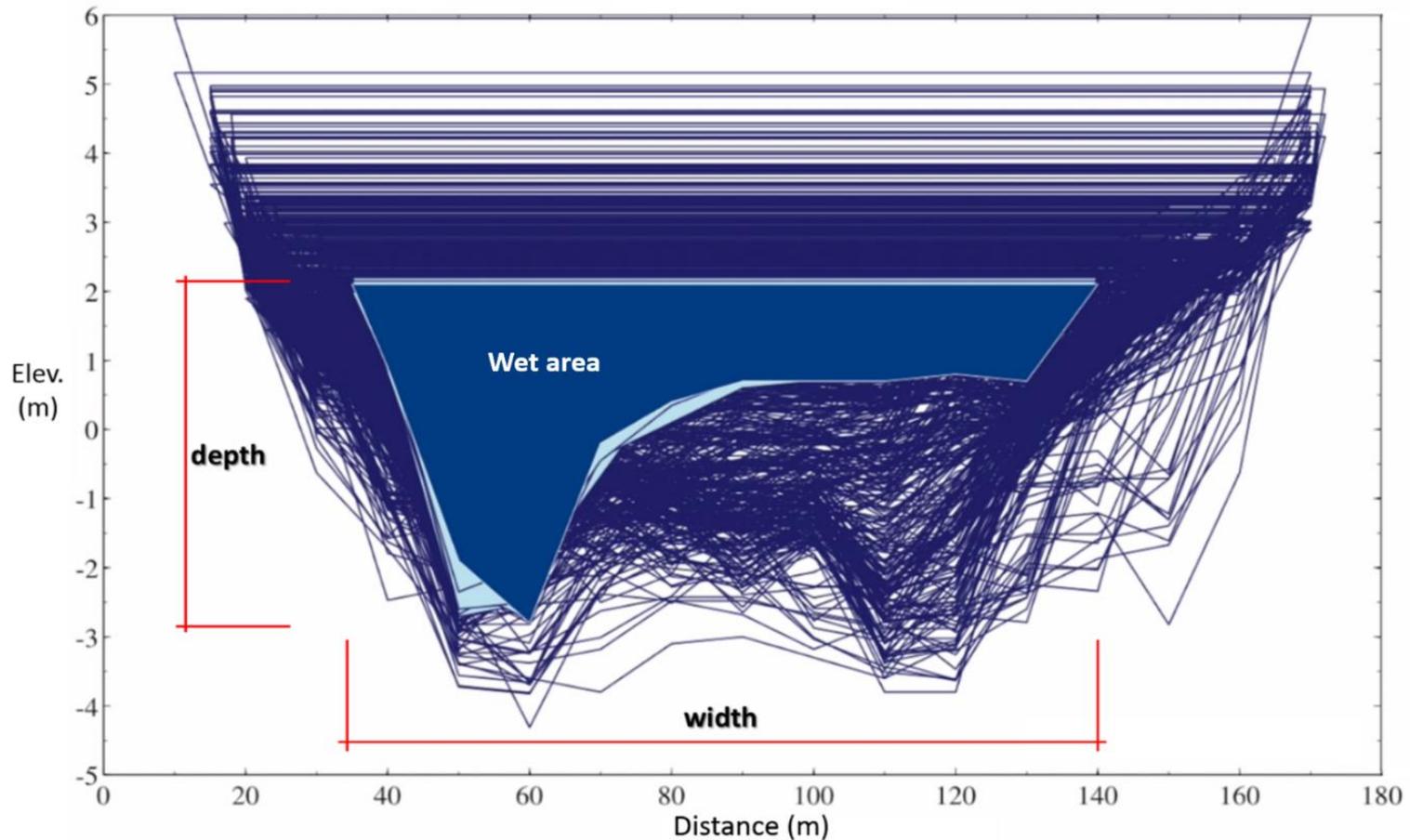
# Effects of discharge and vegetation on channel cross-section - Numerical study Pilcomayo River

(Grissetti, Crosato, Bregoli, 2019)



Strong and quick morphological variations, high sediment transport rates  
Sand-bed, floodplain vegetation, semi-arid climate  
Daily data series on: discharge (up to 4,600 m<sup>3</sup>/s), water levels and cross-sections. Sediment sampling data

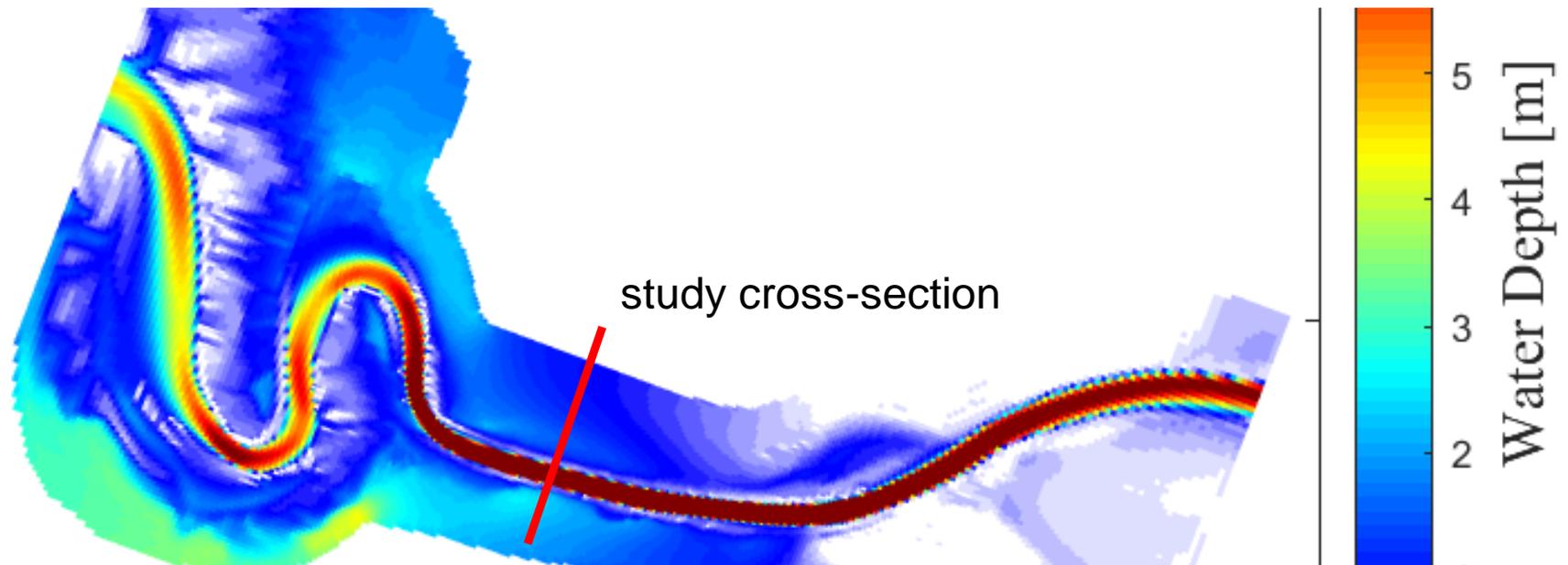
# Strong daily evolution Pilcomayo River: quick adaptation of cross-section to discharge



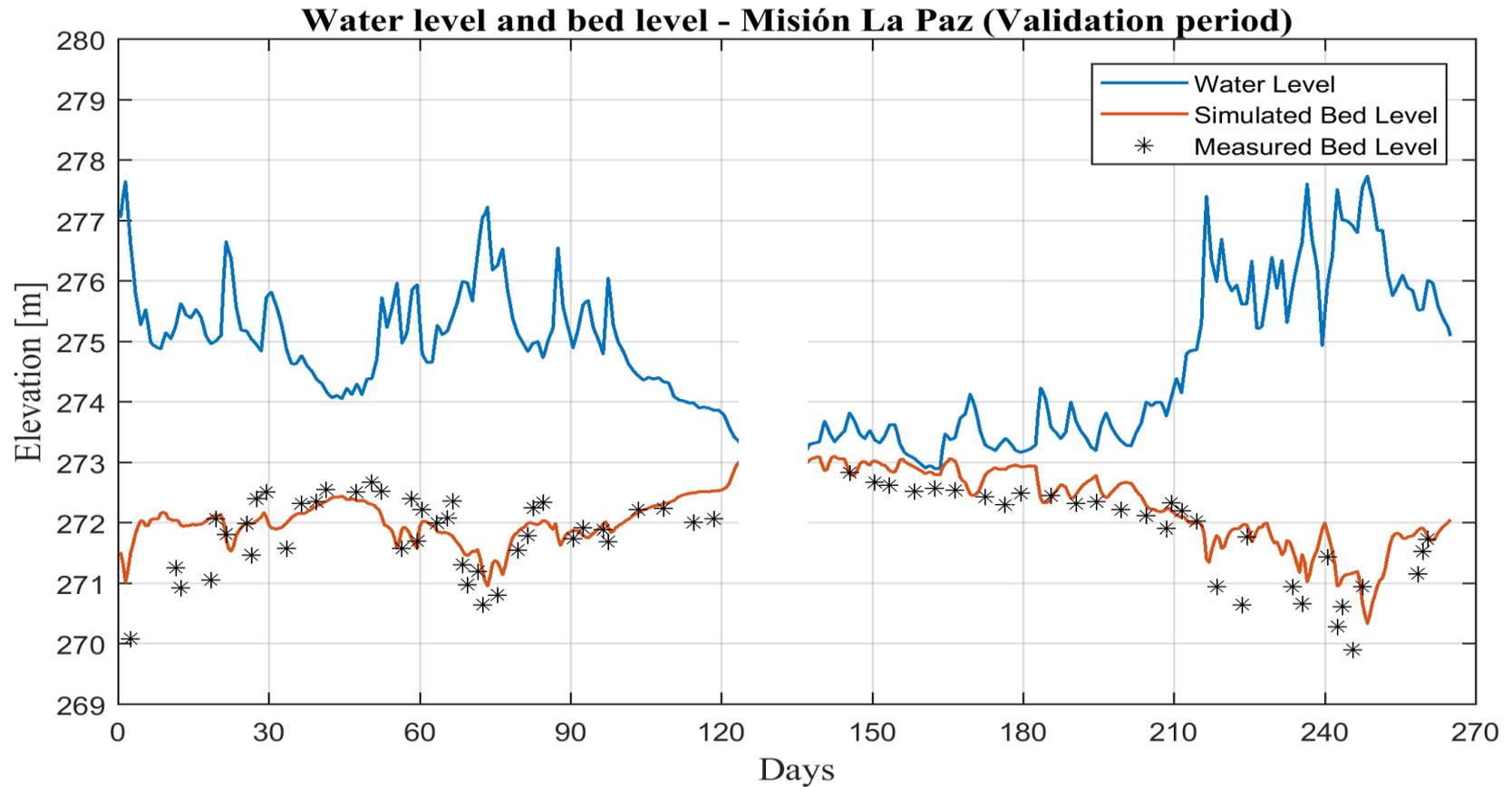
(Capapé and Martín-Vide, 2015)

# 2D Model construction with Delft3D

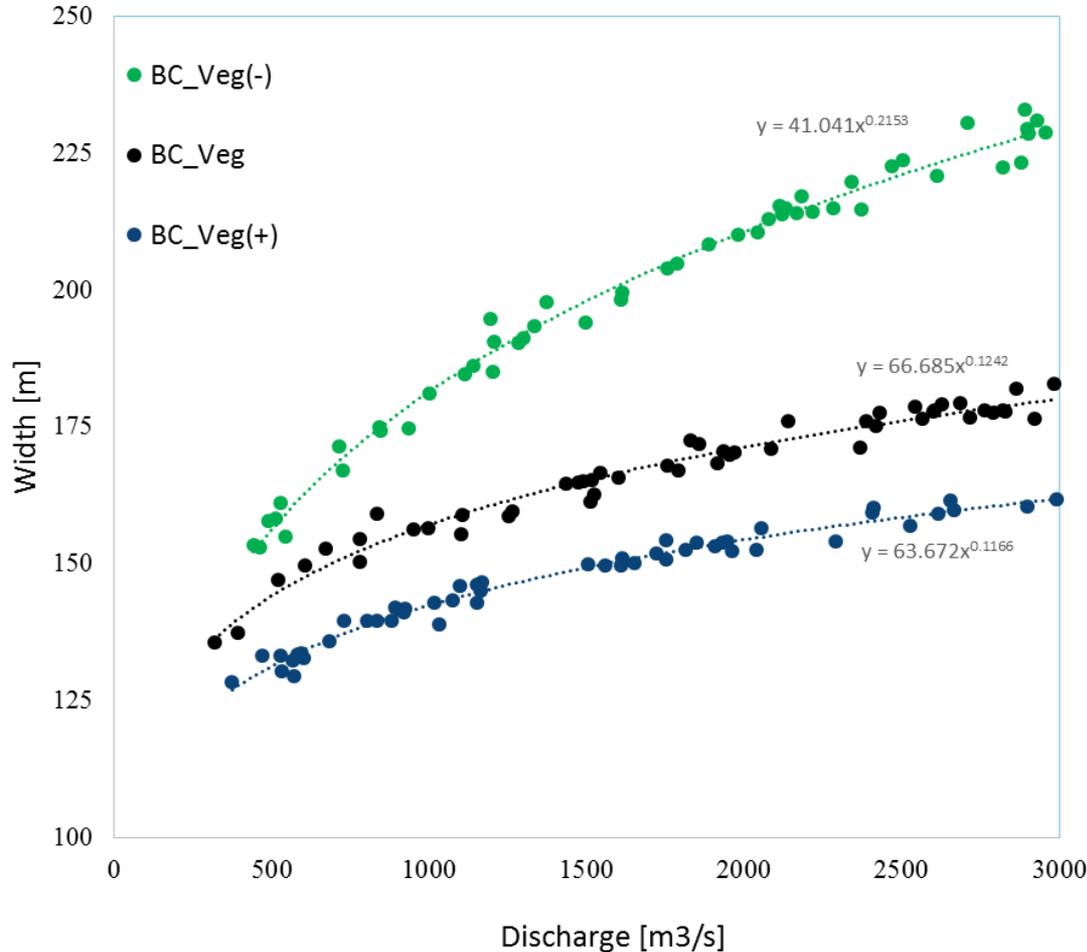
- Model calibration 1972-1973 (Chézy coeff., transverse slope effects)
- Model validation 1980-1981
- Model runs:
  - Present situation (B.C.)
  - Present situation without floodplain vegetation
  - Combinations:
    - More/less floodplain vegetation - Higher/lower discharges



# Results of model validation

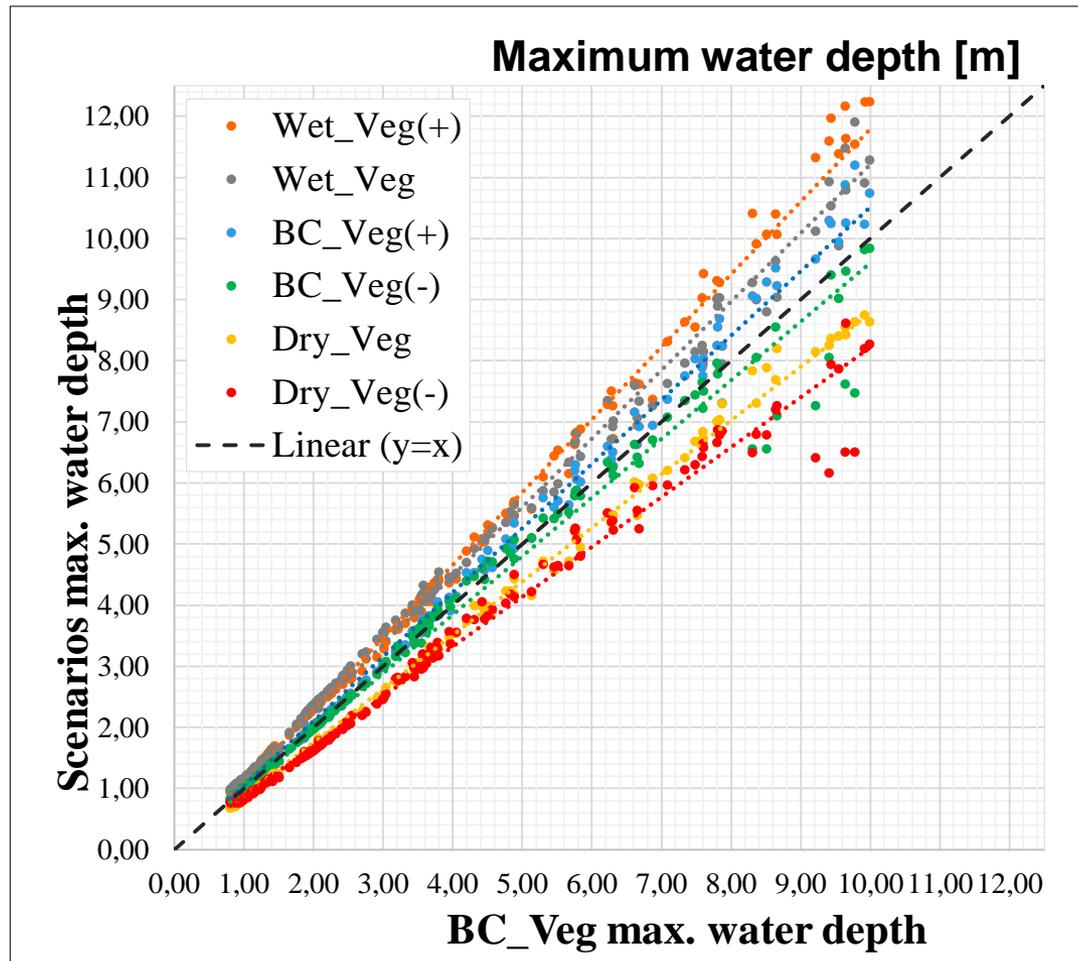


# Results: channel width compared to Base Case



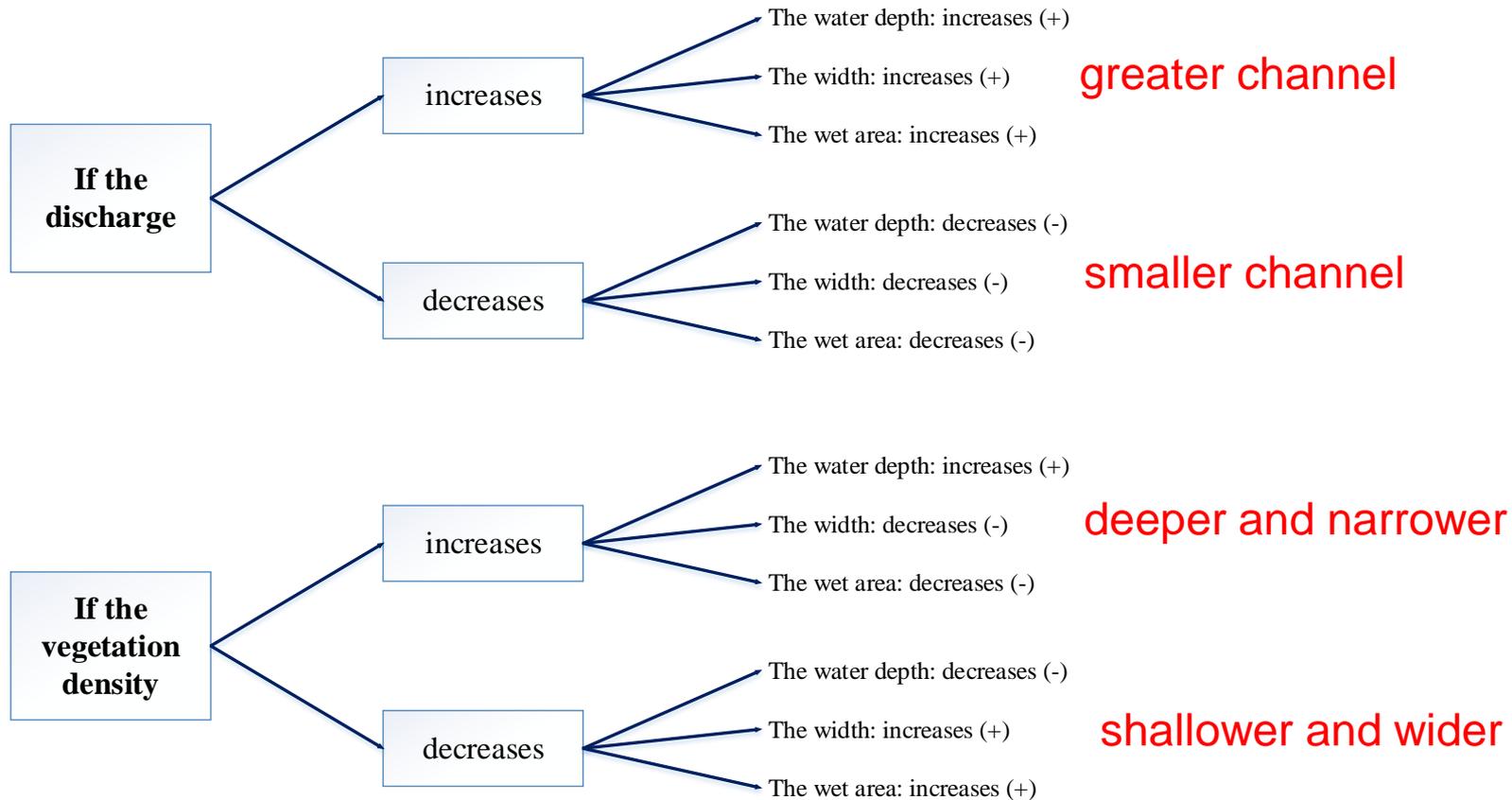
vegetation  
has a strong  
impact on width

# Results: Thalweg depth compared to Base Case



discharge  
has more impact  
on depth  
than vegetation

# River channel response to variations of floodplain vegetation and discharge



**THANK YOU**

**Questions?**

